Accumulation-rate changes through the Holocene from ice-core records and radar data

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Outline

Basics deriving accumulation histories from ice-core records

Ice and accumulation histories from Central WAIS

- Extending ice-core accumulation records spatially with radar
 - Central West Antarctica (shallow layers)
 - Taylor Dome (deep layers)

- Context for recent climate change
 - Antarctic peninsula Holocene temperature history
 - Last 2,000 years of WAIS ice and climate change

Why this matters...

- Ice-core records have high temporal resolution; compare independent records to each other
- Specifics of past climate are a necessary boundary condition to run ice-evolution models for ice-sheet reconstructions
- Radar-detected internal layers are an increasingly abundant data set that can add spatial information
- Holocene climate change gives an interesting and necessary context to modern records

 Deriving climate histories from ice-core records requires knowledge/assumptions about the ice-sheet history

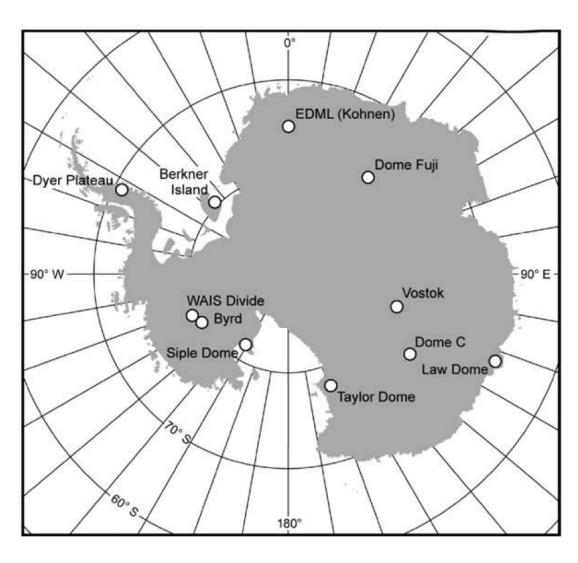
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- Ice-sheet model reconstructions constrained by ice cores
- Temperature histories provided by ice-core data put modern change in context with the past

Deep ice-core sites



- + South Pole
- + Roosevelt Island
- + James Ross Island

Cuffey and Paterson (2010)

Where to drill an ice core ... theoretically

Assume uniform strain rate and no melting

Accumulation constant

In reality, strain rate varies with depth

Cuffey and Paterson (2010)

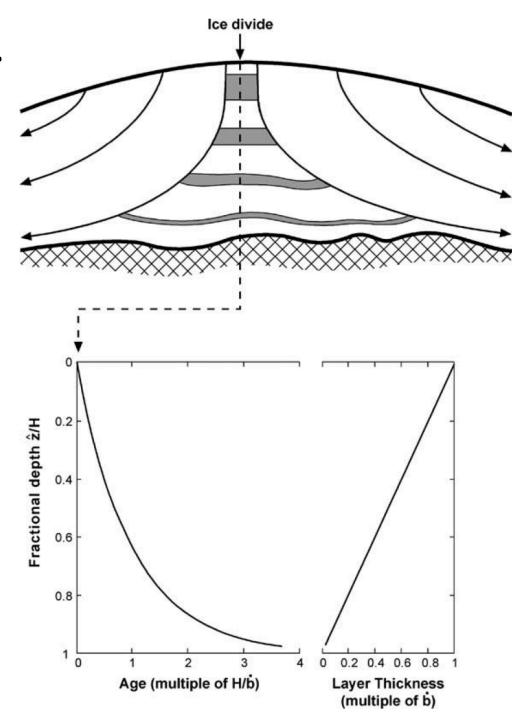


Table 15.1: Ice core ages at 80% of total depth.

| Core Site | Ice thickness (m) | Accum. rate (myr ⁻¹) | Theoretical age (kyr) | Actual age (kyr) | Reference |
|--------------------|----------------------|----------------------------------|--------------------------|---------------------|-----------|
| Agassiz Ice Cap | 111 [†] | 0.098 | 1.8 | 5 | (1) |
| Law Dome | 1220 | 0.7 | 2.8 | 4 | (2) |
| Taylor Dome | 550 | 0.06 | 15 | 50 | (3) |
| GISP2 | 3050 | 0.23 | 21 | 50 | (4) |
| EDML ^{††} | 2774 | 0.07 | 64 | 105 | (5) |
| Dome C | 3275 | 0.027 | 195 | 350 | (6) |

[†] Ice-equivalent value.

References: (1) Fisher et al. 1995; (2) Morgan et al. 1997; (3) Steig et al. 2000; (4) Meese et al. 1997; (5) EPICA Community members 2006; (6) Parrenin et al. (2007).

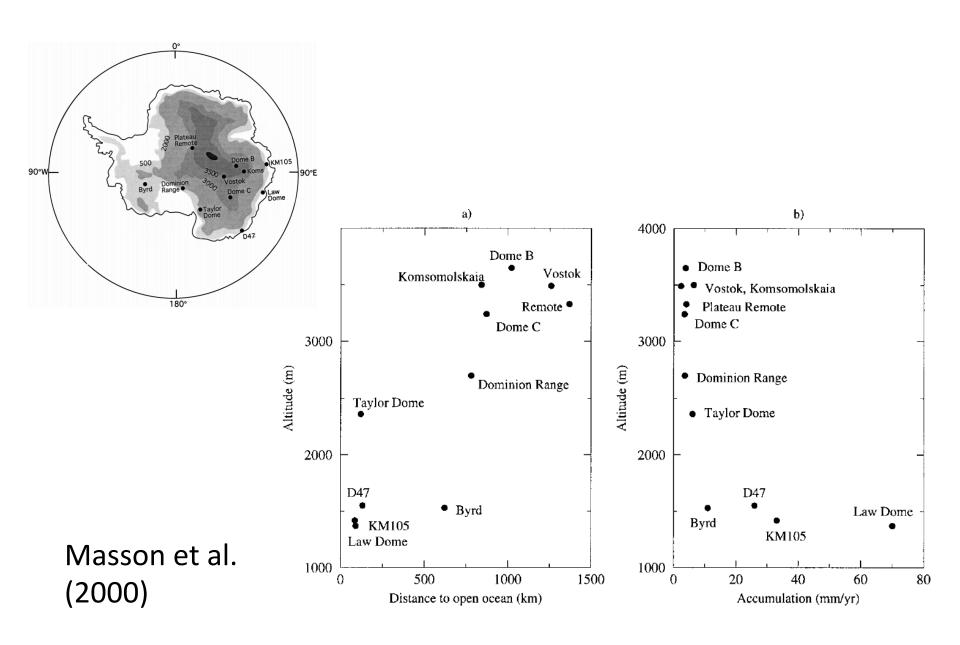
Cuffey and Paterson (2010)

^{††} EDML stands for "EPICA Dronning Maud Land."

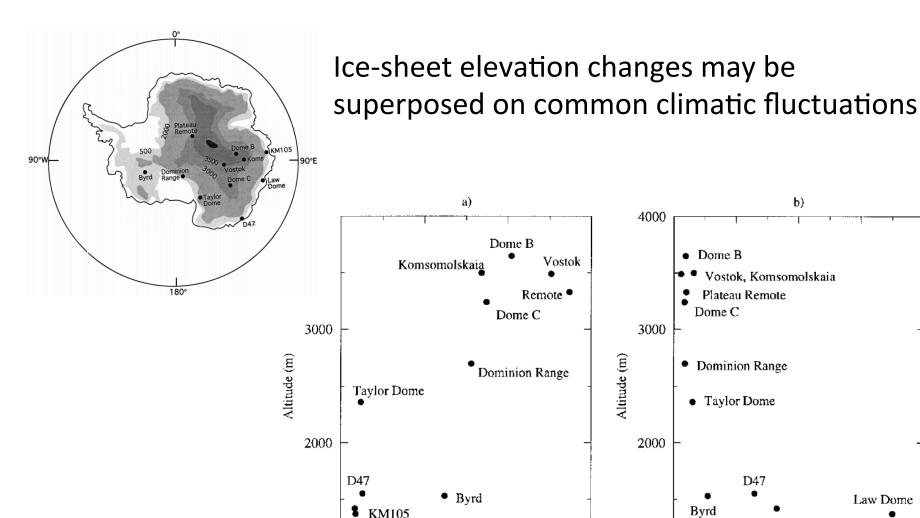
Basics of deriving a climate record: Age scale

- Annual layer counting (chemistry, dust, electrical conductivity)
- Volcanic, cosmogenic, and other marker horizons
- Markers related to globally mixed gases
- Oxygen and hydrogen isotopes of ice
- Physical characteristics (e.g. bubbles)
- Ice-flow model with depth-varying strain rate

Holocene ice-core records across Antarctica



Holocene ice-core records across Antarctica



Law Dome

500

Distance to open ocean (km)

1000

1000

KM105

20

40

Accumulation (mm/yr)

60

80

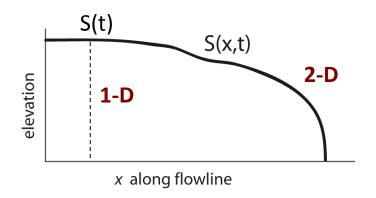
1000

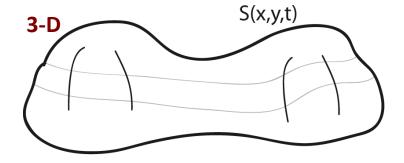
1500

Masson et al. (2000)

Basics of deriving a climate record: Accumulation rate

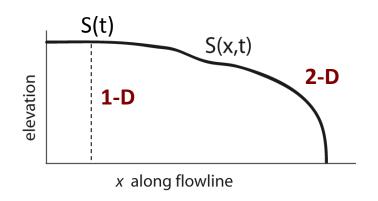
Need to apply **ice-flow models** in order to estimate accumulation history at a core site from measured layer thicknesses

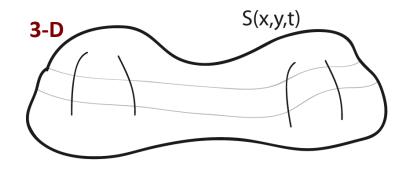




Basics of deriving a climate record: Accumulation rate

Need to apply ice-flow models in order to estimate accumulation history at a core site from measured layer thicknesses





Inferred accumulation rate:

$$\dot{b}_{core}(t) = \frac{\lambda(t)}{\Lambda(t)}$$

Layer thickness -- measured
Thinning function -- calculated

Basics of deriving a climate record: Accumulation rate

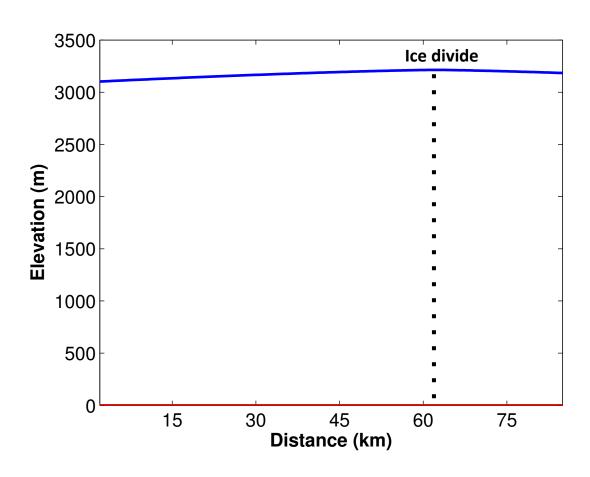
Separate the influence of ice flow vs. climate on ice-core record

Sometimes this is easy ... but, sometimes this is hard...

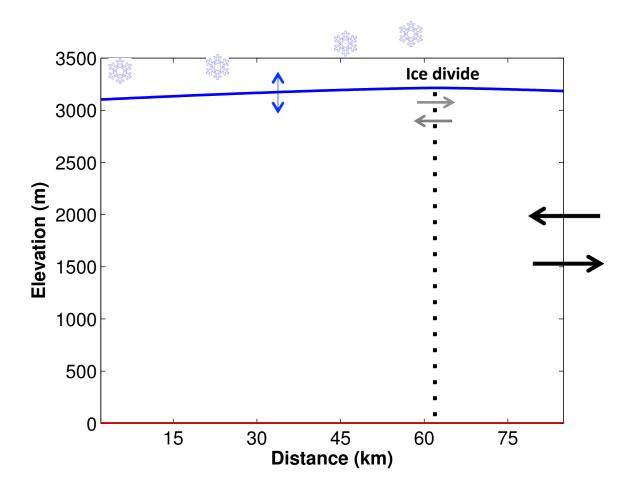
Inferred
$$\dot{b}_{core}(t) = \frac{\lambda(t)}{\Lambda(t)}$$

Layer thickness -- measured Thinning function -- calculated

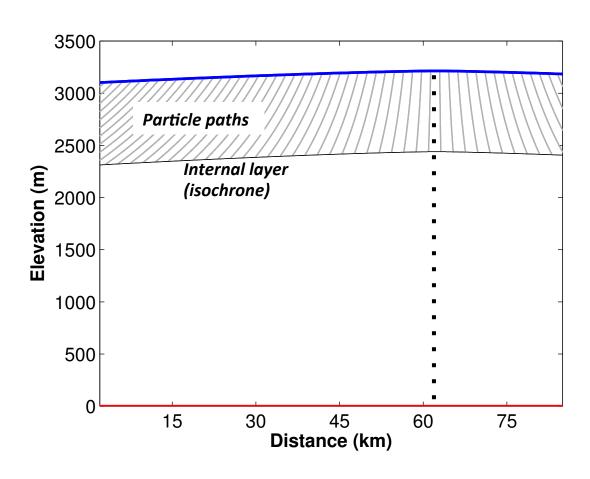
Assume that we can describe ice flow in ice-sheet interior



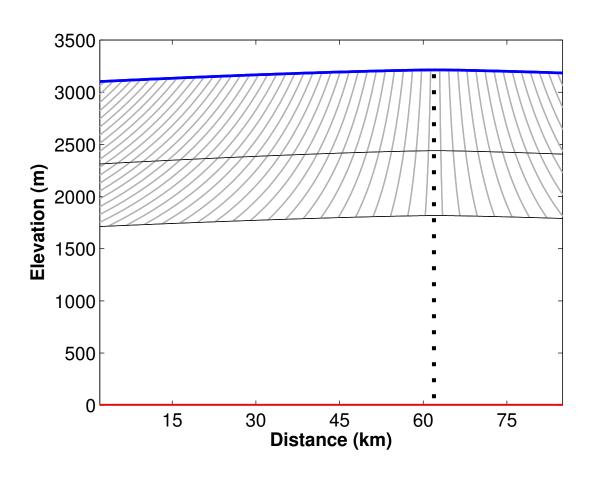
As the ice sheet evolves over thousands of years, accumulation rate and flow rate change



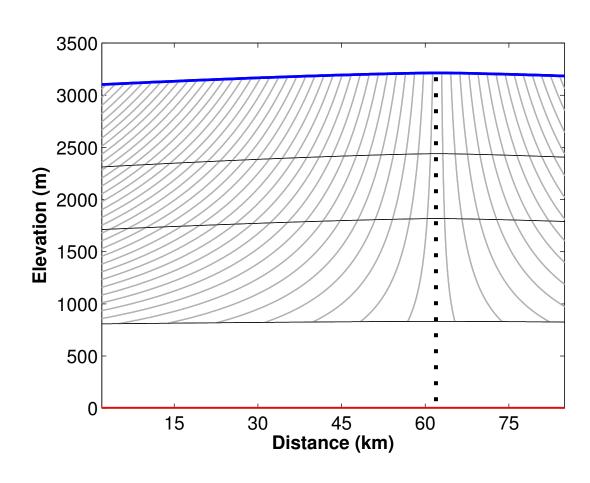
The rate, pattern, and time-variation of climate and ice dynamics affect particle paths



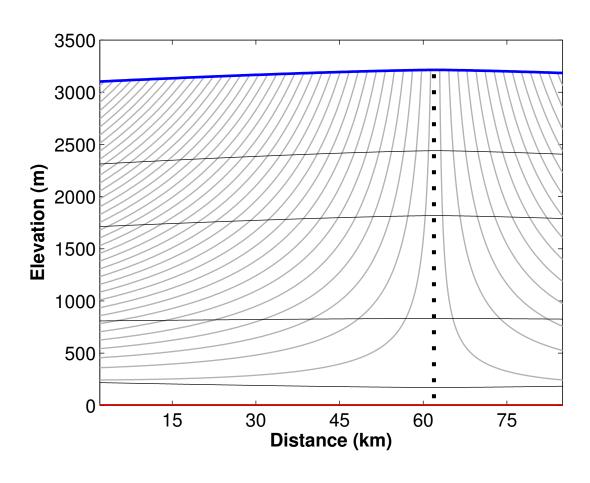
Deeper layers from further back in time



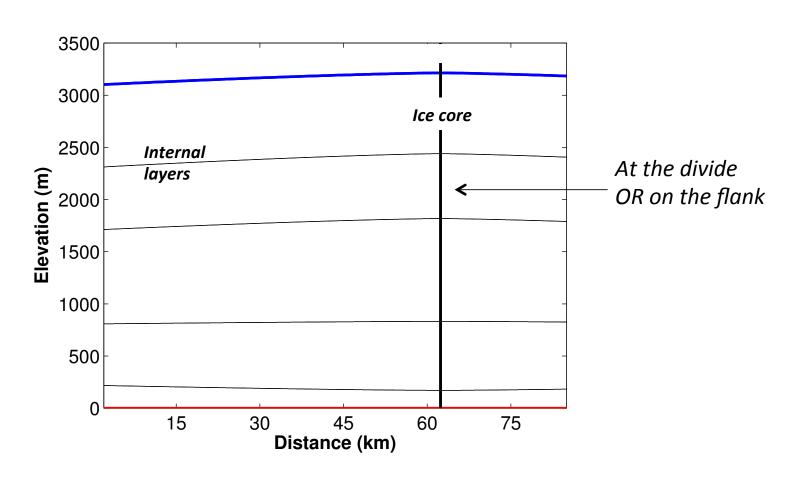
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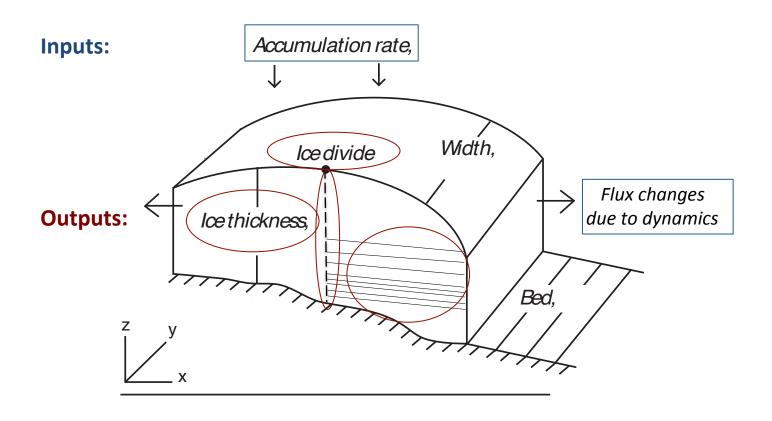


Dynamic and climatic histories recorded in the internal layers and in an ice core:

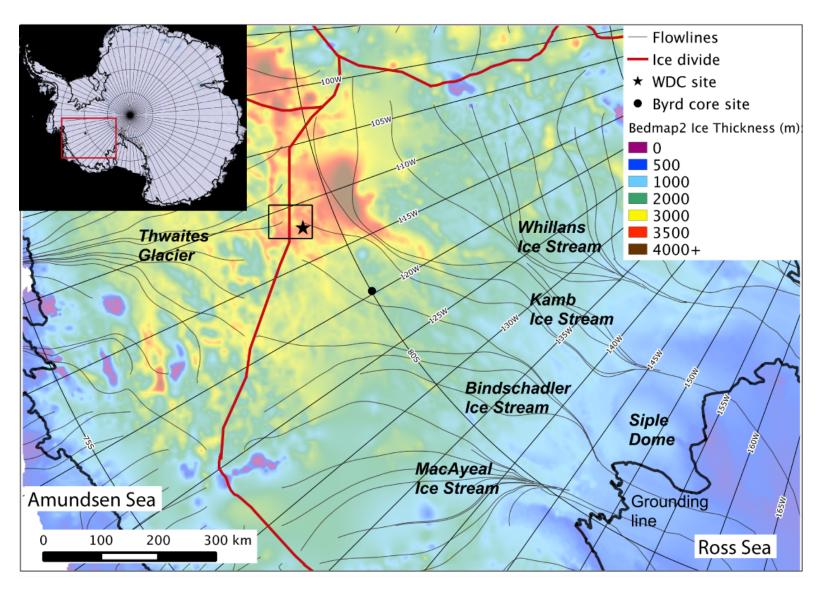


Flowband (ice flow) model

We can use the ice-core record and layers together with an ice-flow model to *constrain* ice-sheet history

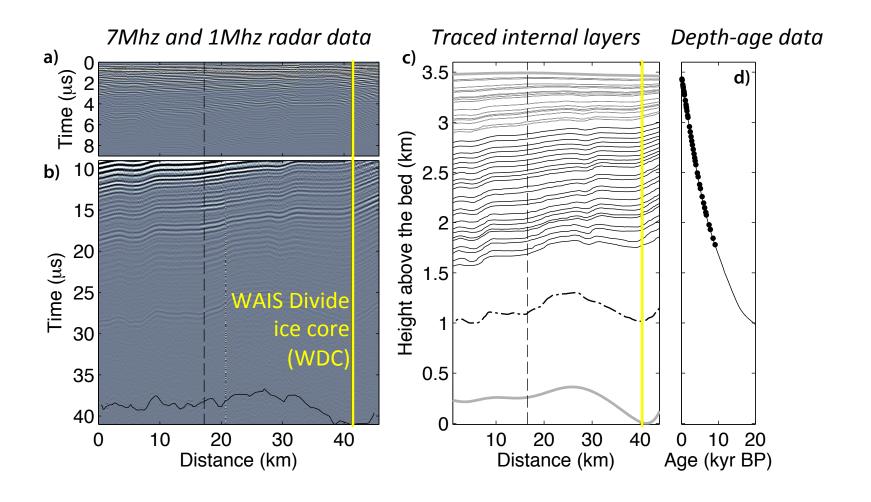


Central West Antarctica

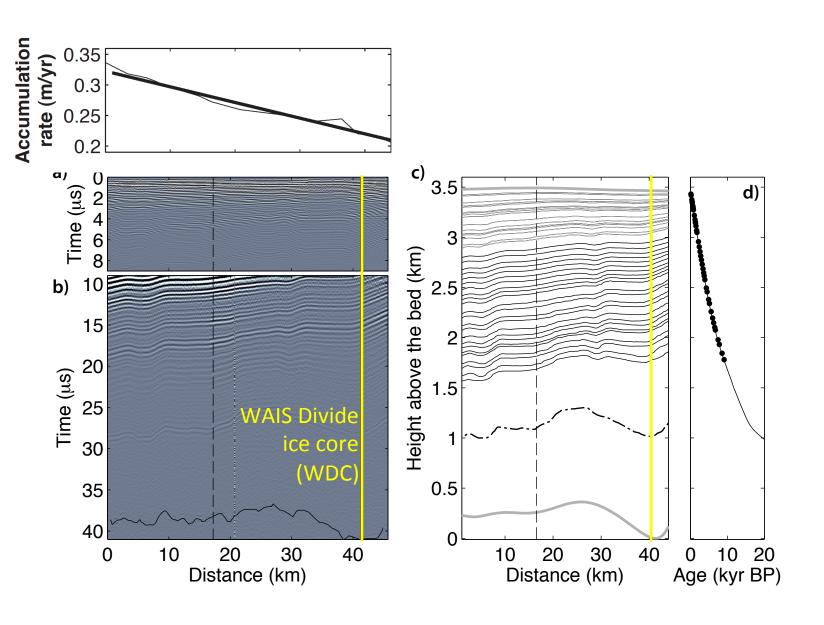


Koutnik et al. (2016)

Central West Antarctica: Data

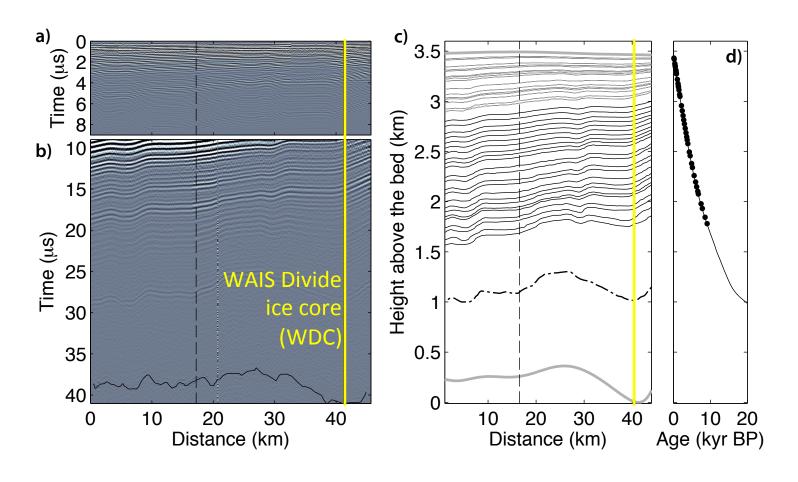


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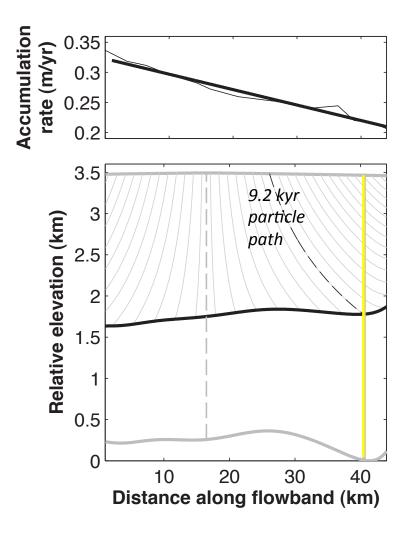


Central West Antarctica: Data

What is the Holocene accumulation history at the ice-core site? Ice divide is migrating today, did it migrate in the past?



Central West Antarctica



Inferring ice and climate histories from the ice-core record must be done carefully...

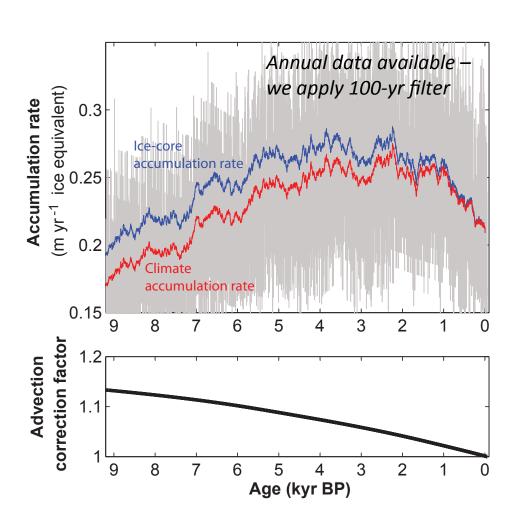
- 1. Ice in core originated upstream
- Apply thinning function from a 1-D model (*Buizert et al.*, 2015)

Then, advection correction to estimate accumulation history ...

Central West Antarctica: Simple estimate of accumulation

Ice-core accumulation rate:

Accumulation at locations where ice in core originated upstream



Central West Antarctica: Simple estimate of accumulation

Ice-core accumulation rate:

Accumulation at locations where ice in core originated upstream

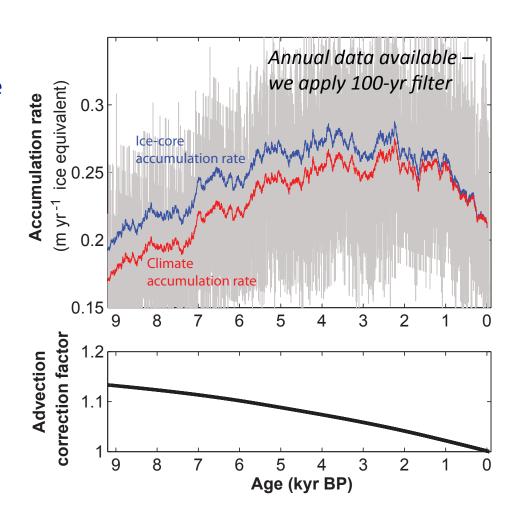
Climate accumulation rate:

Accumulation at fixed location (core site)

Climate accumulation rate =

Ice-core accumulation rate

Advection correction factor



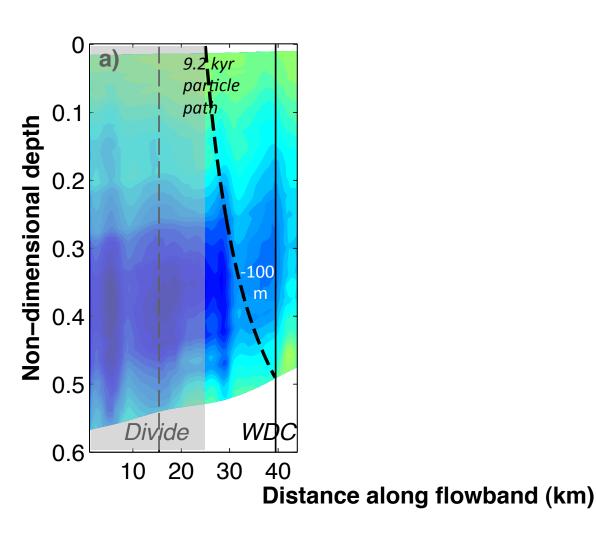
Central West Antarctica

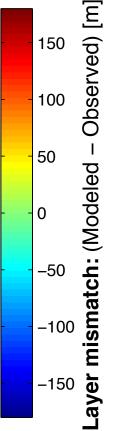
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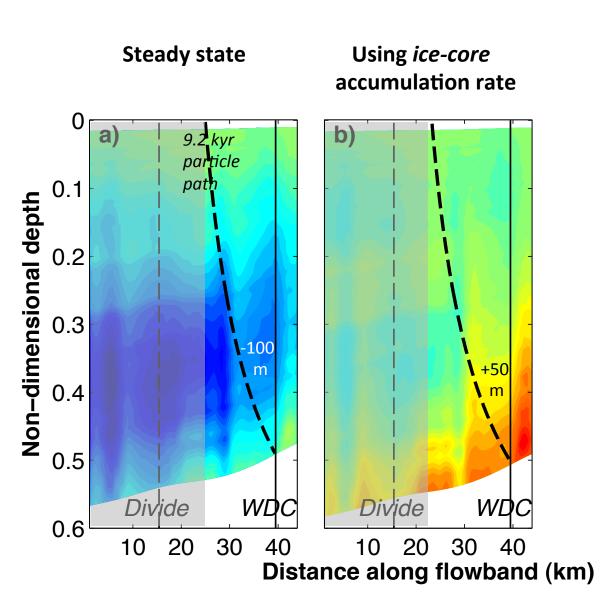
- 1. Ice in core originated upstream
- 2. Apply thinning function from 1-D model and then simple advection correction
- Use climate accumulation history in 2.5-D model to generate layers
- 4. Do modeled layers match observed?

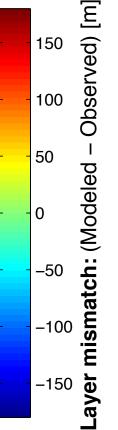
Central West Antarctica: Flowband model calculations

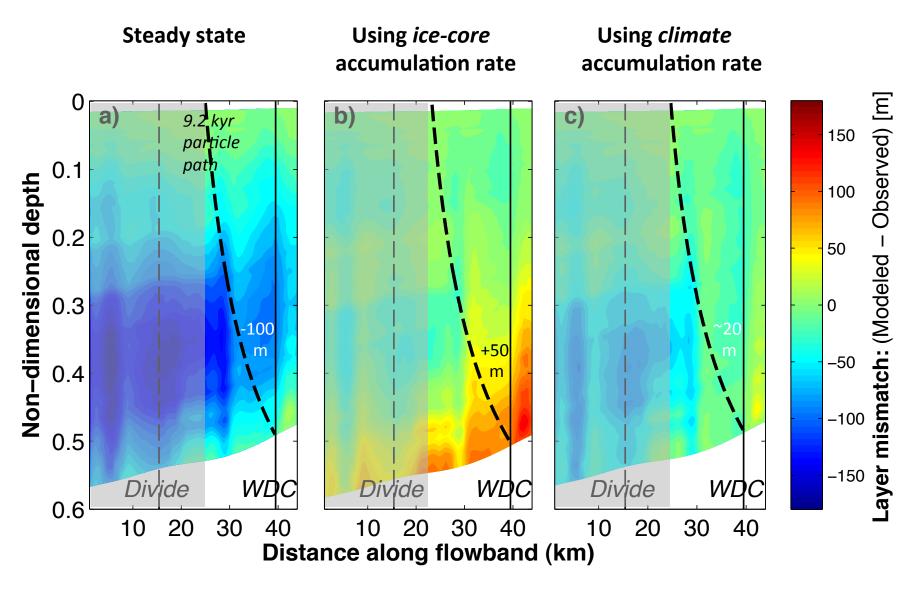
Steady state

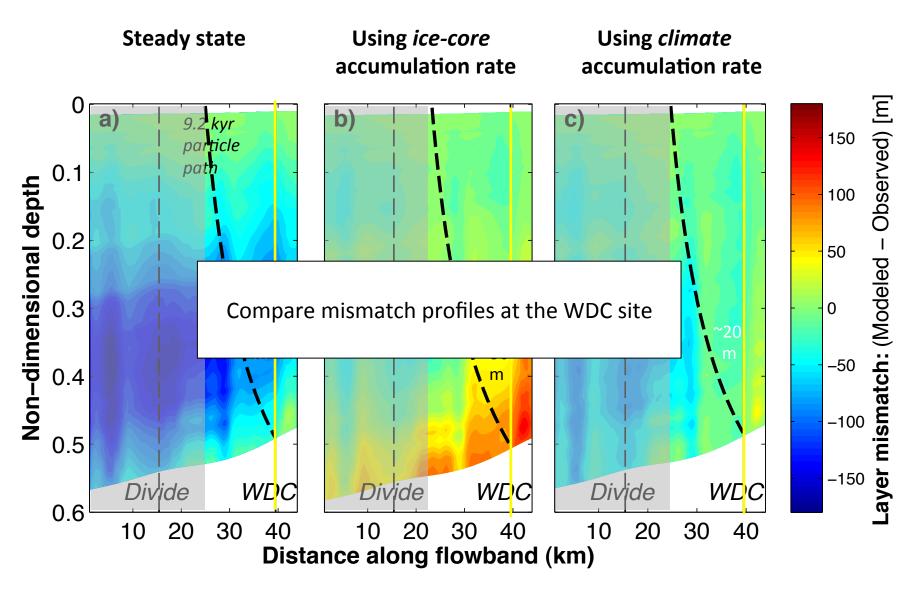








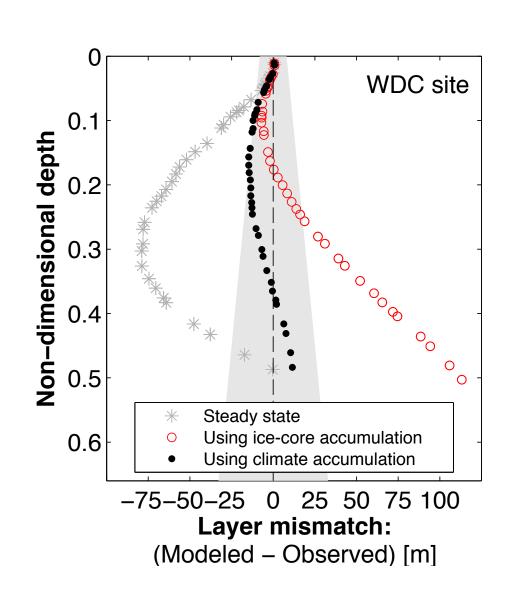




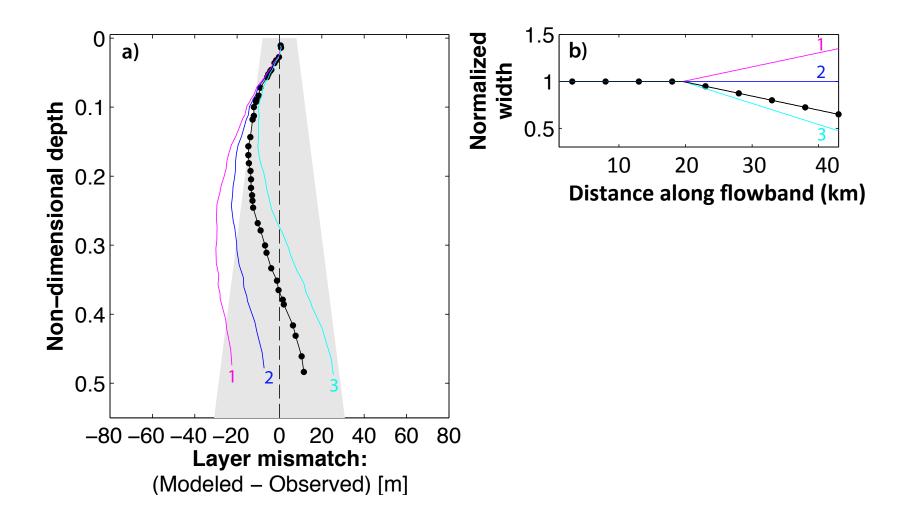
Central West Antarctica: Flowband model calculations

Using climate accumulation history gives a really good fit to internal layers

(within uncertainties in layer depths from data collection, processing, travel-time to depth, and layer picking: ± 8-32 m)

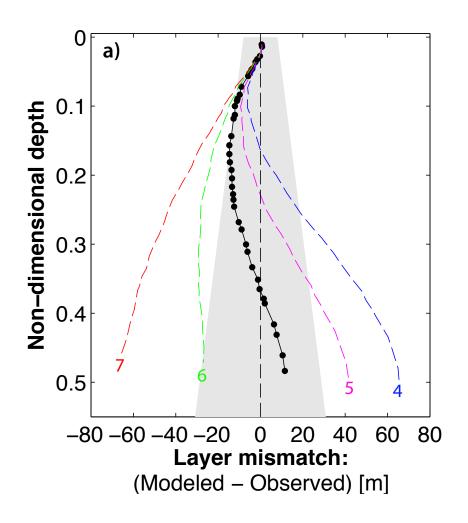


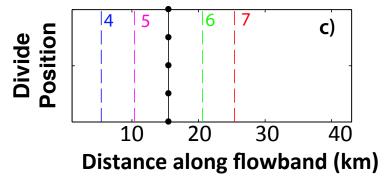
Evaluate the goodness of fit at WDC to poorly known model parameters: Width function



Evaluate the goodness of fit at WDC to poorly known model parameters:

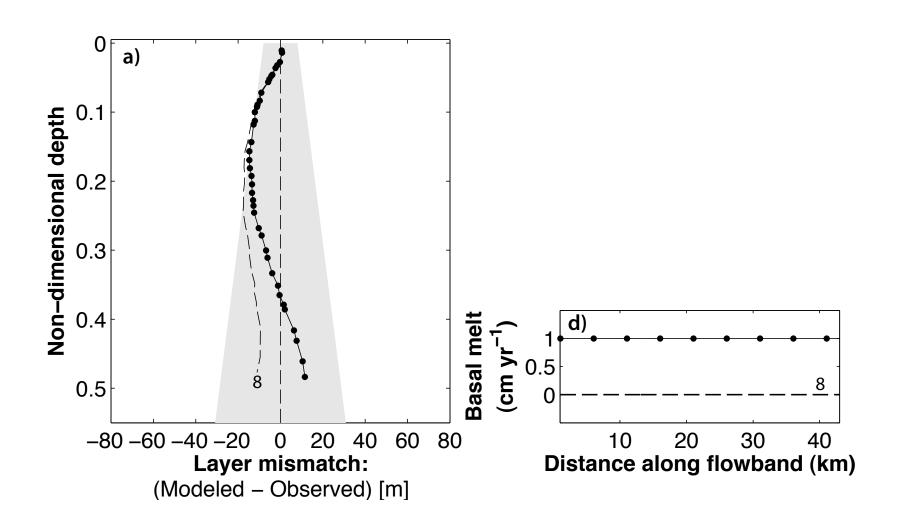
Holocene ice-divide position



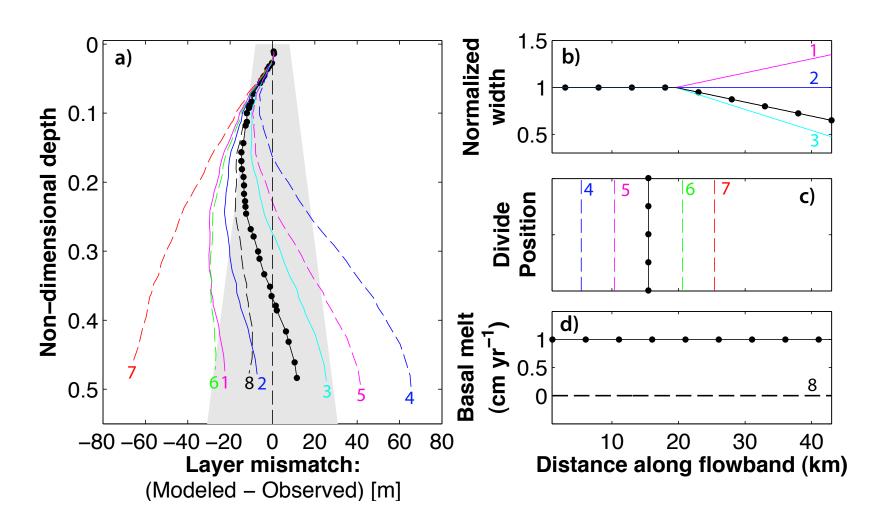


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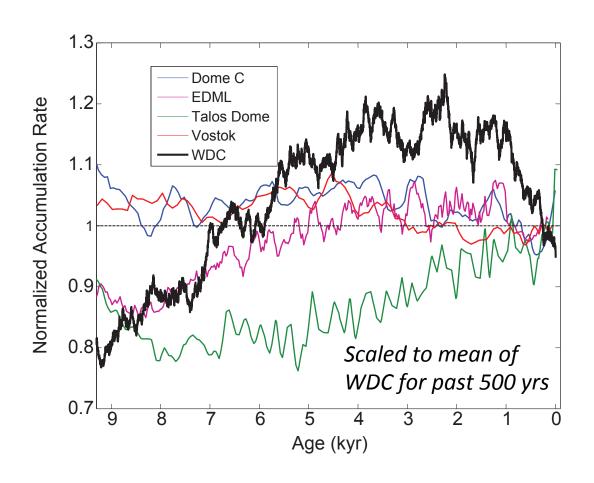
Basal melt rate

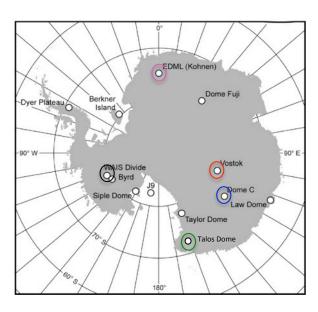


Different width functions and basal melt rates may also be consistent with layers – **different Holocene divide position is not consistent**

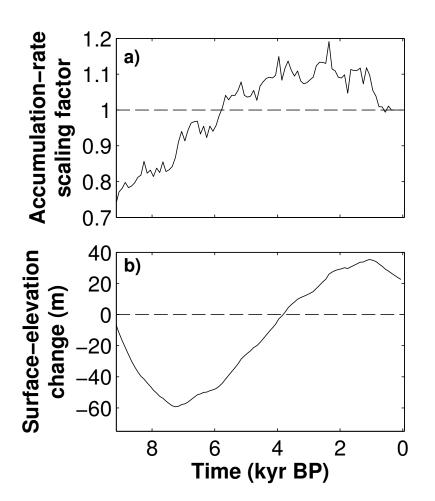


1. Accumulation rate was ~20% lower than present at 9.2 kyr BP, and there was ~40% increase from 9.2 - 4 kyr BP \rightarrow big change



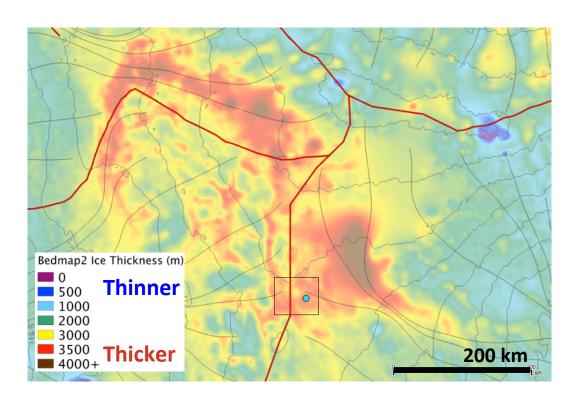


2. Change in accumulation leads to change in interior ice thickness unless compensated by a change in ice dynamics



** Calculated elevation change depends on initial condition

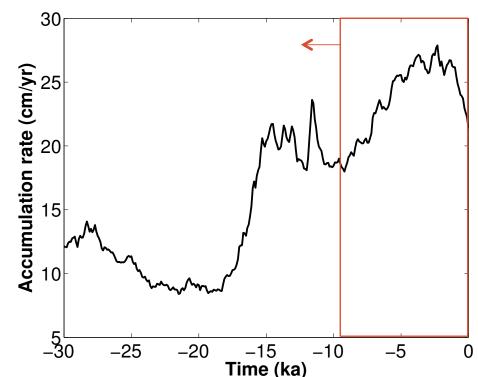
3. The average Holocene divide position likely remained within ~5 km of the modern position



The bedrock topography may have macroscopic control on ice thickness and divide position

Annual layer-counted timescale allowed for the first independent accumulation reconstruction back to 31 ka





Fudge et al. (2016)

Central West Antarctica: Implications

1. To infer the most realistic accumulation history from the ice-core record, need to correct for advection of ice from upstream

2. Climate models extrapolate West Antarctic climate from East Antarctic ice-core records – we can now do better

3. Continent-scale model reconstructions should use this climate forcing, and calculated Holocene changes in ice geometry should include a relatively stable divide near WDC

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What do continent-scale reconstructions show?